

Recent (1995–1998) Canadian research on contemporary processes of river erosion and sedimentation, and river mechanics

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Abstract

Canadian research on contemporary erosion and sedimentation processes covers a wide range of scales, processes, approaches and environmental problems. This review of recent research focuses on the themes of sediment yield, land-use impact, sediment transport, bed material transport and river morphology and numerical modelling of fluvial landscape development.

Research on sediment yield and denudation has confirmed that Canadian rivers are often dominated by riparian sediment sources. Studies of the effects of forestry on erosion, in-stream sedimentation and habitat are prominent, including major field experimental studies in coastal and central British Columbia. Studies of sediment transport mechanisms have focused on the composition of particles and the dynamics of sedimentation. In fluvial dynamics there have been important contributions to problems of turbulence, scale, flow structure and entrainment processes, and the characteristics of bedload transport in gravel-bed rivers. Although

addressing broader environmental concerns[However\ large reductions in programme funding for water resources in Canada in the 0880s "Pearce and Quinn\ 0885# have had severe consequences for long!term sediment monitoring programmes[Monitoring is now reduced to about 199 stations\ most of which are in Ontario\ Alberta and British Columbia\ and only about 09) of stations have a full regular monitoring programme[The remaining 89) are either based on occasional sampling to establish rating curves "in Alberta# or use monitoring only during spring freshet or major storm events "data and information based on personal communication from Environment Canada personnel P[Zyrmiak\ B[Smith and T[Yuzyk\ March 0888#

Previous studies\ based mainly on Federal Government sediment load data\ have shown that sediment loads in the glaciated landscape of Canada are often dominated by riparian sources\ with uplands con!tributing little to the sediment load and landscape denudation "Ashmore and Day\ 0877^Church e a [\ 0878^ Ashmore\ 0882# This pattern of erosion can also be shown by using sediment particle characteristics rather than sediment yields and budgets "De Boer and Crosby\ 0885# For example\ sediment sources in the Assiniboine Whitesand River system in Saskatchewan are de!coupled so that sediment generated in the headwater portion of the basin is not transported downstream\ but instead is deposited within the system[The sediment carried in the downstream reaches is derived locally from the areas directly adjacent to the mainstream "De Boer and Crosby\ 0885# Analysis by SEM:EDS has been used to demonstrate that suspended solids in the upstream portions of the basin consist predominantly of planktonic diatoms\ with small numbers of composite particles comprised of mineral grains cemented together with clay and organic matter[In the downstream portions of the basin\ suspended solids consist of composite particles cemented together with clay minerals\ and few diatoms are present "Crosby and De Boer\ 0885# The downstream change in suspended solids characteristics is explained by the increased contribution to the sediment load of reworked glacial\ glacio~uvial and glaciolacustrine deposits on the main valley ~oor[

Where uplands are erosionally connected to the stream system\ the e}ect of di}ering sur_cial "glacial# deposits on sediment yield may be detected[Stone and Saunderson "0885# investigated the regional patterns of ~uvial sediment yield in the low relief\ largely agricultural\ tributaries to the lower Great Lakes using Water Survey of Canada data for 81 tributaries[The average speci_c sediment yield of the drainage basins of the Great Lakes and Lake St Clair ranged from 10=6 to 76=2 t km⁻¹ year⁻¹\ but the highest speci_c sediment yields " 499 t km⁻¹ year⁻¹# were found in sub!basins with agricultural and industrial land use on _ne!grained glacio!lacustrine parent materials[However\ these results are di. cult to reconcile with soil loss observations on the Prairies "Pennock e a [\ 0884# showing that the highest rates of soil loss were found in glacial till landscapes "29 t ha⁻¹ year⁻¹ 2999 t km⁻¹ year⁻¹# whereas the lowest rates were found in silty glaciolacustrine and aeolian landscapes and in _ne sandy loam glacio~uvial lacustrine landscapes "01 t ha⁻¹ year⁻¹ 0199 t km

year

example recent research in southern Manitoba "Hargrave and Shaykewich\ 0886# shows that the bulk of nutrient "N\ P# losses are associated with sediment rather than soil water[The timing of loss is also important and there is increasing evidence that snow melt runoff\ rather than rainfall\ is the major agent of erosion in many parts of Canada "McConkey *e a* \ 0886^ Edwards *e a* \ 0887^ Gill *e a* \ 0887# This is consistent with the high sediment concentrations observed during spring runoff\ events in streams[However\ high erosion rates vary with land treatment and high erosion rates also may be accompanied by high rates of redeposition within the _eld "Edwards *e a* \ 0887# The redeposition process also calls into question the reliability of soil loss equations calibrated on very short "a few metres# erosion plots[Evidence from sheet wash experiments "Mathier and Roy\ 0885# indicates that the parameters of sediment transport in sheet wash equations for agricultural _elds are applicable over slopes only between 3 and 44 m[

E \ *ed e a a d ea ab a e ec e a ve g*

Fluvial processes have an important effect on in-stream biological habitats[The effects of timber harvesting on erosion and in-stream sedimentation are particularly prominent in research on watersheds in coastal British Columbia "Hogan *e a* \ 0887a# but the concerns extend to many other areas of the country[

In British Columbia attention now has been focused also on interior areas through an interdisciplinary study of six experimental watersheds in the Stuart Takla region[The basins are in the Hagem Range of the Omineca mountains at the northern end of the sub-boreal spruce biogeoclimatic zone forming the most northern extent of the Fraser River basin[The streams support both early and late run sockeye salmon "O'Connell *e a* # as well as other species of _sh "Macdonald and Herunter\ 0887# One of the major goals of the project is to determine how forestry-associated changes in the thermal\ hydrological and geomorphological regimes in these watersheds affect variation in the structure\ stability and distribution of _sh habitat "Tschaplinski\ 0887# Baseline "pre-harvesting# studies of snow hydrology\ groundwater influences and road-related sediment sources were undertaken for 4 years prior to any harvesting activity "Heinonen\ 0887# Suspended sediment loads in three of the streams have been monitored since 0881 "Beaudry\ 0887# and indicate that the pre-harvested watersheds have very low annual suspended sediment yields compared with other British Columbia watersheds[Approximately 84% of the annual yield is delivered in a 10 month period during snowmelt "Beaudry\ 0887# Hogan *e a* \ "0887b# have mapped the channel morphology of three streams in early autumn for each year of the study to evaluate changes associated with hydrological regime and large woody debris positioning[Pre-harvest data show large variability in the longitudinal profiles of these natural streams[If harvesting causes the long profile to become less variable\ as suggested in coastal watershed studies\ the effects will be detectable statistically "Hogan *e a* \ 0887b# Abundant natural large woody debris is present in all the channels and is generally very stable[Hogan *e a* \ "0887b# indicate that the source of the woody debris is the upstream reaches\ which necessitates protection of both the _sh habitat in the lower reaches and the woody debris sources upstream[

There are similar forestry concerns in eastern Canada[The Copper Lake buffer zone study\ in the boreal forest near Cornerbrook\ Newfoundland\ has been assessing the impact of forest harvesting on brook trout "Sawyer *a* # populations and habitat since 0882 "Scruton *e a* \ 0884# Clarke and Scruton "0886# evaluated the _ne!sediment yields in two streams upstream and downstream of road construction using the Wesche method of sediment trapping "partially buried sediment box traps filled with gravels# for evaluating statistical differences in _ne!sediment storage[Seasonal differences in sediment storage\ and also spatial

The absence of any apparent impact of logging roads at other sites may be explained by a lack of significant hydrological events between construction and sampling

TRANSPORT AND FLOCCULATION OF FINE SEDIMENT] PROCESSES AND IMPACTS

Fine sediment processes are the dominant agents of physical erosion in watersheds and an important transfer mechanism of contaminants. These processes are a complex interaction of physical, chemical and biological processes in streams and much of the recent work has focused on the details of these mechanisms, especially the formation and destruction of flocs, channel morphology and the effects of flocculation on sediment transport.

Flocculation and erosion

It is well established that fine sediment in streams is transported as flocs rather than as individual particles, yet there is much to be learned about the formation and structure of flocs and the implications for sediment

Fractal dimensions may be used to quantify changes in the morphology of rocks. De Boer (1986) Images of rocks from a small prairie stream were analysed and fractal dimensions of the rock population were determined from the relationships between rock area, perimeter and length. The changes in the morphology of the rocks resulting from algal bloom could be quantified with the fractal dimension derived from the area-perimeter relationship. The change in fractal dimension from 0.15 to 0.31 represents the formation of large particles with intricate shapes and large perimeters during the algal bloom. Fractal dimensions were also used by De Boer and Stone (1987) to describe the effect of buffer strips on rock morphology in two contrasting drainage basins in southern Ontario prior to and during snowmelt. In the Strawberry Creek basin, the narrow, ineffective riparian buffer zones resulted in an influx of sediment-laden overland flow from farmland. As a result, during snowmelt, sediment concentrations increased by a factor of 29 compared with pre-melt conditions, and an increase in the fractal dimension from 0.13 to 0.24 reflected the change in rock shape resulting from the input of fines and organic matter from topsoil. Conversely, in the Cedar Creek basin the extensive riparian buffer zones prevented an influx of particles during snowmelt, resulting in little change in sediment concentrations and fractal dimensions.

Contaminant transport

It is now recognized that much contaminant transport in rivers is associated with fine-grained sediments. In Canada this has led to both basic research and management initiatives on major river systems such as the Fraser, the St Lawrence and the Peace and Athabasca (NRBS, 1985; Gray and Tuominen, 1987) with a focus on anthropogenic effects and the impact of sediment-borne contaminants on the entire ecosystem (Ouellet, 1986; Cary and Cordiero, 1986).

It is apparent that the significance of sediment-borne contaminants depends on both the nature of the contaminants and the in-stream sediments. Stone et al. (1984) showed that phosphate release from fine-grained bed sediment depends on both the sorption characteristics of the individual particle size fractions and the particle size distribution as a whole. Different contaminants also show differing size fractions. Stone and Droppo (1985) For example, in southern Ontario streams, concentrations of Zn and Pb were greatest in the smallest $7\ \mu\text{m}$ fraction, whereas Cu concentrations were greatest in the 7 to 01 $\mu\text{m}</math> fraction. Overall, 79% of the load of Pb, Cu and Zn is associated with particles $20\ \mu\text{m}$, which illustrates the important role of fines in metal transport. Trace-metal studies in the St Lawrence River (Quemerais et al., 1985) showed that Co, Mn and Fe are transported predominantly in the particulate phase, whereas a significant proportion of Cd and Cu (45 and 37%, respectively) is transported in the dissolved phase. Quemerais and Lum (1986) report that particular Cd concentrations in the St Lawrence River showed a positive correlation with particular organic matter and Mn, and a negative correlation with suspended solids concentration. The latter was explained by the influx of relatively coarse suspended sediment, with a low ability to adsorb trace elements, by the Quebec tributaries during high discharge events.$

Transfer of heavy metals on street sand is an important concern in urban areas. Stone and Marsalek (1985) investigated the trace element composition and partitioning of street sediment collected in Sault Ste Marie, Ontario and found that a large portion of the heavy metals are potentially bioavailable. The elevated levels of Cd, Pb, Cu, Zn, Mn and Cr in the exchangeable and/or soluble phases indicate that street sediment has a potentially adverse effect on water quality during runoff and snowmelt. Nevertheless, a large proportion of the total metal load was associated with coarser particles, which have little potential for downstream transport.

TURBULENT FLOW STRUCTURE, SEDIMENT TRANSPORT AND BEDFORMS

Introduction

Increasing interest in the study of small-scale processes in rivers parallels an international trend apparent since the 1970s. For example, recent symposium volumes include Clifford et al. (1982), Ashworth et al. (1985), Carling and Dawson (1985). Further impetus for the study of turbulence in sedimentology and

geomorphology has come from the discovery of quasi-periodic structure in boundary layer flows and their potential role in the development of alluvial bedforms (Leeder, 1987). Small-scale flow structures are centrally implicated in sediment transport (Church, 1985), although it is debatable whether short-lived (a few seconds) structures are of sufficient duration to directly affect the development of fluvial landforms. Overall turbulent characteristics of the flow may be more relevant to fluvial geomorphology than are individual flow structures. Recent work has focused on the overall changes in flow turbulence produced by specific bed morphology, on the detection and characterization of intermittent flow structures both in space and time, and on the detection and description of the particle-scale bed structures in gravel-bed streams. These studies cover small-scale flow processes in a wide range of contexts (bedrock channels, gravel-bed rivers) and morphologies (river confluence, pebble clusters) with a dominant interest in gravel-bed rivers. Only a few have successfully addressed the issue of fluvial landform development as a consequence of the interactions between flow turbulence and sediment transport.

Bed structure in gravel-bed rivers

Particle clusters are a common mode of particle organization on gravel-bed streams, but more recently Church *et al.* (1987) have shown that larger scale organization and particle-size segregation may occur. Their field observations are corroborated by laboratory study. The most common structures are stone cells that

Flow structure and bed discordance

Because of their critical role in the routing of flow and sediment, river confluences have been the subject of much research in the last decade. Recently, it has become apparent that flow structure at confluences is strongly affected by the "common" discordance of the conluent beds. Laboratory experiments have clearly shown the effects of bed discordance for parallel channels (Best and Roy, 1988) and the structures are now known for confluences with a 29° angle (Biron et al., 1985a,b). One of the main effects of bed discordance is distortion of the mixing layer between the two flows. The width of the mixing layer increases rapidly near the apex of the confluence and some fluid from the deeper channel is entrained under that of the shallower channel in such a way that upwelling occurs near the bank on the side of the shallower channel. The entrainment of fluid in the lee of the step is probably due to the low pressure zone that develops there. As a result, the separation zone usually occurs on the side of the tributary as it enters the confluence and is much smaller when the beds are concordant than when they are discordant. Field study in a sand-bed confluence also shows that bed discordance affects the three-dimensional flow structure and the intensity of the turbulence generated in the shear layer, and distorts the mixing layer (De Serres et al., 1988). The back-to-back secondary flow cells often reported at symmetrical confluences were not observed at this site. These results are beginning to be corroborated by three-dimensional numerical flow models (Bradbrook et al., 1987). Although this work is directed primarily towards understanding the physical processes of river confluences, it has important environmental applications. For example, bed discordance enhances mixing intensity and reduces the mixing length downstream of confluences (Gaudet and Roy, 1984), indicating that bed morphology is a primary factor to consider in the environmental management of river confluences.

Turbulence and bed adhesion

Suspension of sandy bed material is also strongly controlled by turbulent flow structures. Direct and simultaneous measurement of turbulent eddies "using electro magnetic flow meters" and local suspended sediment concentration "using optical backscatter sensors" make it possible to study the intermittent nature of suspension (events) and the influence of mean flow conditions (Lapointe, 1985). In a river about 0.99 m wide and 1.4 m deep, turbulent momentum exchange is highly variable, and exchange levels 0.9 times the mean occur up to 4% of the time but account for up to 69% of the stress. Vertical sediment flux 0.9 times the mean rate occurs up to 4% of the time, but accounts for up to 89% of the net suspension and is more frequent in deeper flows, perhaps related to dune development. The overall frequency distributions are similar to those found in much larger rivers, suggesting the occurrence of a general frequency distribution for vertical mixing. Suspension appears to be dominated by eddies with length scales one to five times flow depth.

Dunes in the Fraser River estuary channels (Kostaschuk and Villard, 1985) are of two types: symmetrical with rounded crests and low angles (< 7°) and asymmetrical with steep lee side (> 0.8). An important feature of these dunes is the absence of lee-side flow separation, although it may be present intermittently in the lee of asymmetrical dunes. Round-crested dunes are characterized by higher sediment transport rates than the asymmetrical dunes because of higher near-bed flow velocity. Asymmetrical dunes represent a transitional morphology between the large symmetrical dunes and smaller dunes. The low transport rate on their stoss side is caused by the presence of smaller dunes superimposed on the larger form, and lead to lower deposition rates on the lee side. The Fraser estuary dunes fit in the upper-stage plane bed region of

The critical state can be maintained only by substantial increases in flow resistance which Tinkler (1986a) argues comes mainly from intense vorticity and shearing between the zone of critical flow and the surrounding flow. The initiation of bedload transport may also contribute to the increase in roughness and initial motion of boulders may take place beneath standing waves. Carling and Tinkler (1987) Turbulent flow processes and structures are also important for fluvial erosion of bedrock and the micro morphology of bedrock channels. Erosion marks such as grooves and ripples often found in bedrock channels are caused by the mechanical action of suspended sediment and the most intense erosion is associated with flow separation eddies around protrusions and large vortices developed in the mixing layer between high velocity layers and stagnation zones. Tinkler (1986b) and Hancock *et al.* (1987)

BED LOAD TRANSPORT, SEDIMENT SORTING AND CHANNEL MORPHOLOGY

Although bedload is often a very small component of the total sediment load of most rivers, it is critical to the development of river morphology. Recent Canadian research has tackled three major issues related to the processes of bedload transport in gravel-bed streams: the actual downstream changes in particle size in relation to terrain and tributary characteristics; the relationship between river morphology and bedload transport processes; and description of the particle population of which bedload is composed. The first issue is completely addressed in the book by Carling and Tinkler (1987) and the second in the book by Tinkler and Carling (1987). The third issue is addressed in the book by Carling and Tinkler (1987) and the fourth in the book by Carling and Tinkler (1987). The book by Carling and Tinkler (1987) is a comprehensive review of the current state of knowledge on bedload transport in gravel-bed streams. The book by Tinkler and Carling (1987) is a comprehensive review of the current state of knowledge on the relationship between river morphology and bedload transport processes. The book by Carling and Tinkler (1987) is a comprehensive review of the current state of knowledge on the description of the particle population of which bedload is composed. The book by Carling and Tinkler (1987) is a comprehensive review of the current state of knowledge on the relationship between river morphology and bedload transport processes. The book by Tinkler and Carling (1987) is a comprehensive review of the current state of knowledge on the description of the particle population of which bedload is composed.

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channel change and bed material transport rates[For example\ on the Vedder River\ British Columbia "Martin and Church\ 0884# it is apparent that depositional zones "downstream decrease in transport rate# are wider and more braided than the single!channel {transport zones|[This reach!scale pattern is persistent over several years\ although there are some changes in reach behaviour over time and depending on ~ood magnitudes[

are slightly less precise than scour chains for estimating active area\ but are preferred for practical reasons *
the tracer particles are already needed for virtual velocity estimates[

RETROSPECT AND PROSPECT

The characteristics of the Canadian landscape and environment have always conditioned and guided ~uvial

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